



GEOS 657 – MICROWAVE REMOTE SENSING
GRADUATE-LEVEL COURSE AT THE UNIVERSITY OF ALASKA FAIRBANKS

Lecturer:
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Lecture 12: InSAR for Deformation Monitoring, Automatic Processing Services & Limitations of Traditional InSAR



UAF Course GEOS 657

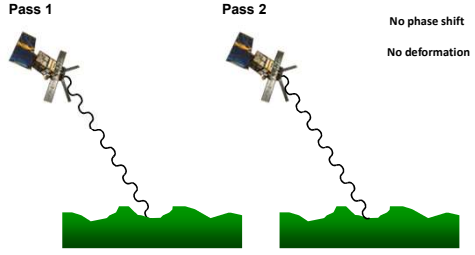


- <https://avo.alaska.edu/images/image.php?id=194280>
- <https://avo.alaska.edu/images/image.php?id=194281>




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How InSAR Works



- InSAR phase is a function of distance from satellite to ground (range)

Pass 1 Pass 2 No phase shift
 No deformation



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InSAR Processing – Steps 1 – 7:
Find and Download Images, Co-Register & Form Interferogram

• To form interferogram, we calculate: $I = u_1 \cdot u_2^*$ (where $*$ is complex conjugate)

Example: Mt. Peulik volcano, Alaska

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InSAR Processing – Step 8:
Subtraction of Flat Earth Phase

• Example:
– ERS-2 Interferogram of Mt. Peulik volcano, Alaska

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DIFFERENTIAL IN SAR – A METHOD FOR CM-SCALE DEFORMATION MONITORING

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The Concept of Differential InSAR (d-InSAR)

Interferometric Phase:
 $\phi = \phi_{\text{topo}}(z; B) + \phi_{\text{defo}}$
 $\phi_{\text{defo}} = \frac{4\pi}{\lambda} \Delta R_{\text{defo}}$

d-InSAR Goal: extraction of deformation signal from interferometric phase

terrain motion or subsidence

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How to Separate Topographic and Deformation Phase?

- If reliable DEM is available, use DEM to compensate for topography.
- Else, ≥ 3 complex SAR images at times t_1, t_2, \dots, t_n are required:
 - Form several interferograms:
 - time lag: $\Delta t_{n-m} = t_n - t_m$
 - baseline: $B_{\perp, n-m}$
 - phase: ϕ_{n-m}
 - For constant velocity: $\phi_{n-m}^{\text{defo}} = \phi_{n-m} \propto \Delta t_{n-m}$ and $\phi_{n-m}^{\text{topo}} = \phi_{n-m} \propto B_{\perp, n-m}$
- For singular displacement event: use $\phi_{n-m} \propto B_{\perp, n-m}$ to derive topography

For signal separation consider:

- Deformation phase changes only with time
- Topography phase changes only with Baseline

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Example of Topography Compensation

- Compensation using available DEM
- ALOS PALSAR data of Baja Earthquake, 2010

1 Fringe: 12 cm motion along Line-of-sight

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How well can we measure deformation?

How to measure surface motion from the InSAR phase: $\phi_{defo} = \frac{4\pi}{\lambda} \Delta R_{defo}$

ΔR is motion in sensor look direction
 $\Delta R = \Delta y \sin \theta - \Delta z \cos \theta$

For previous PALSAR example:
 1 fringe (2π) corresponds to
 12.5 cm in R
 14.5 cm in z (e.g. subsidence)
 25.0 cm in y (motion)

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Monitoring Volcanoes with InSAR

Principle

- InSAR is capable of observing inflation and deflation of volcanoes
- Inflation and deflation is triggered by changes of magma pressure in magma chamber (see animation)
- These phenomena precede volcanic eruptions and are potentially interesting for predicting eruptive behavior

Westdahl Volcano Deformation Movie

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Interferometric Phase Interpretation

Example: Surface Deformation at Mt. Peulik, Alaska

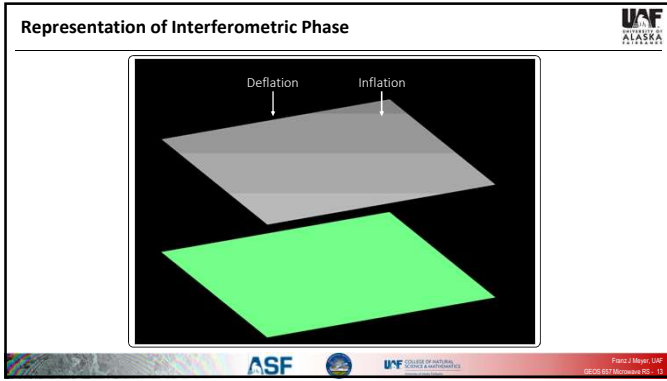
17-cm inflation, Sept. 1996 to Oct. 1997

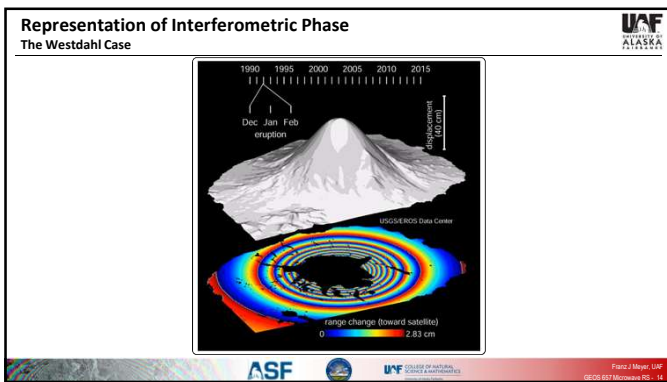
Mount Peulik volcano, Becharof Lake, Ugashik caldera

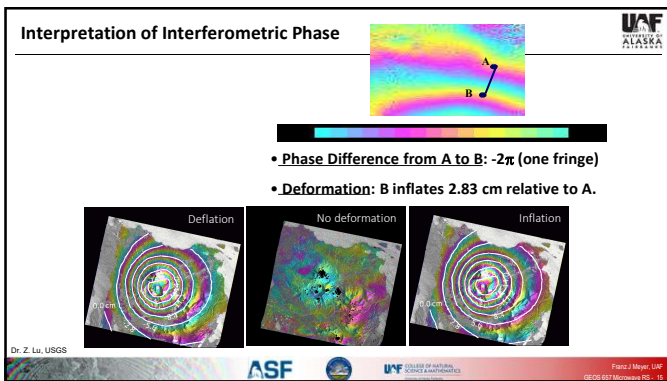
Easting (km)

Zhong Lu, et al. (2001)

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Interferograms in Support of Emergency Management

Example: 2023 M7.8 Turkey–Syria Earthquake

Modern SAR Sensors such as Sentinel-1 and the upcoming NISAR mission can provide InSAR data within 2-3 days of an event

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Continent-Wide Deformation Mapping from InSAR

Example: Antarctica

- First Antarctic-wide glacier velocity map in history
- Full coverage by merging data from a wide range of satellite systems
- Accuracy varies with number of multiple coverage per area and coherence

Mouginot et al., "Mapping of Ice Motion in Antarctica Using Synthetic Aperture Radar Data", Remote Sensing, 4(5), pp 2753-2767, 2012

Courtesy: B. Scheuchl, UC Irvine

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Sub-Milimeter Surface Analysis


Building Deformation in Berlin, Germany

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 (c) 2011 S. Genshardt TUM, DLR

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Typical Applications of Differential InSAR

Application	Notes
• Infrastructure (buildings; roads; levees) deformation	X-band
• Sink holes	X-band
• Landslides	X- or C-band
• Oil/gas extraction	
• Glacier motion mapping	
• Volcano monitoring	
• Permafrost-related surface deformation	C- or L-band
• Soil moisture variation	L-band
• Co-seismic earthquake deformation	L-band
• Post & inter-seismic deformation	L-band
• Ice sheet monitoring	C- or L-band



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Think – Pair – Share


Picking the right data for your InSAR analysis:

- Okmok is a volcano in Alaska’s Aleutian Chain.
- Since its last eruption in 2008, the volcano has been re-inflating at a rate of ~10cm/year

Activity: Pick the most suitable Interferogram for measuring this deformation. Motivate your answer!

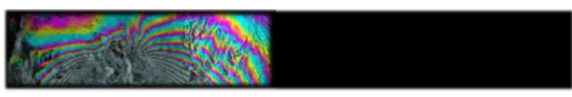
- L-band data; $\Delta t = 24$ days; June - July
- L-band data; $\Delta t = 48$ days; June - August
- L-band data; $\Delta t = 144$ days; July - December
- C-band data; $\Delta t = 24$ days; June - July
- C-band data; $\Delta t = 48$ days; June - August
- C-band data; $\Delta t = 144$ days; July - December

• Please consider the expected coherence γ and phase noise σ_ϕ in your answer



Okmok Volcano, Alaska

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TYPICAL INSAR PROCESSING WORKFLOW

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A Typical InSAR Processing Workflow

1. Select and order InSAR-capable SAR data from a data server
2. Import SAR data into an InSAR processing system
3. Calculate spatial baseline & apply spectral (wavenumber) shift filtering (not discussed in the lectures – applied automatically by most available tools)
4. Determine co-registration parameters:
 - cross-correlate >100 image chips spread over image
 - use over-sampling and interpolation to locate correlation peaks
 - apply regression to parameterize co-registration (e.g. affine transform)
5. Co-register images:
 - Resample slave image(s) to match master image
 - Required accuracy: << 1/10 resolution element
 - More on required co-registration accuracy on Slides 20 & 21

A Typical InSAR Processing Workflow

6. [Optional Orbit Improvement: if precise orbit information is available.]
7. Interferogram formation: $I = u_1 \cdot u_2^*$; optional multi-looking may be applied.
8. Flat Earth phase removal: Simulate and subtract phase trend due to the geometry changes from near range to far range.
9. Coherence Calculation: Coherence is calculated as described in Lecture 12.
10. For differential InSAR (d-InSAR): Using a DEM, simulate and subtract interferogram replicating topography-related phase.
11. Apply phase filter: A phase filter is applied to reduce InSAR phase noise and reduce phase unwrapping complexity (see next section).
12. Phase Unwrapping: Turns originally ambiguous interferometric phase into unambiguous absolute phase.
13. Geocoding and Terrain Correction: Note that flat earth phase needs to be added before geocoding to obtain absolute phase.

Effects of Coregistration Errors on Coherence

• How accurately must the two SAR images be co-registered before interferogram formation?

overlapping area: "signal"

non-overlapping area: "noise" (decorrelation)

ground resolution element of pixel $[i, k]$ in SAR image #1

ground resolution element of pixel $[i, k]$ in SAR image #2

mis-registration in range

mis-registration in azimuth

Stripmap rule-of-thumb: 1/10 of a resolution element mis-registration is usually acceptable

Coherence Reduction Caused by Image Mis-Registration

- Here a plot how coherence drops with pixel misregistration:
 - Assumption: Mis-registration is only cause of de-correlation
 - Mis-registration expressed in fractions of pixels

Coherence γ

Pixel Mis-Registration [%]

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Tips for Selecting Suitable Images for InSAR

How to select a suitable image pair for successful InSAR processing

- Required conditions:**
 - Images from identical orbit direction (both ascending or both descending)
 - Images with identical incidence angle and beam mode
 - Images with identical resolution and wavelength (usually: same sensor)
 - Images with same viewing geometry (same track/frame combination)
- Recommended conditions:**
 - For topographic mapping: Limited time separation between images (temporal baseline)
 - For deformation mapping: Limited spatial separation of acquisition locations (spatial baseline)
 - Images from similar seasons / growth / weather conditions

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AUTOMATIC (INSAR) PROCESSING SERVICES

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Why Is There a Need For Automatic SAR Processing Services?

- **Processing Flows for Generating Value-Added Products Are Often Complicated**
 - Difficult and heterogeneous data types and data formats
 - SAR data typically do not come geocoded and have strong geometric distortions
 - Strong noise effects require complex filtering approaches
 - InSAR processing is complicated and error prone
- **Large Data Volume of SAR Requires Powerful Processing Machines**
- **SAR Processing Software is Often Not User Friendly**


→ Several automatic processing tools are in development to make SAR data more accessible especially to new users from the disaster monitoring communities



Search and Process (In)SAR Data with Vertex




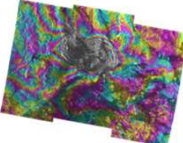
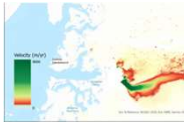
- **Search & discover** ASF's holdings over your AOI
- Use our baseline and SBAS tools to **configure your perfect InSAR stack**
- Submit Sentinel-1 scenes for **on-demand processing** to RTC or InSAR
- **Download** data for additional processing using convenient interface options




HyP3 On-Demand Processing

HyP3: On-demand processing service for analysis-ready SAR data

- Available to anyone at no direct cost
- DAAC managed L1 → L2 workflows


<p>RTC</p> <ul style="list-style-type: none"> • Radiometric Terrain Correction • Corrects geometric and radiometric distortions caused by the terrain 	<p>InSAR</p> <ul style="list-style-type: none"> • SAR Interferometry • The phase measurements of 2 SAR images are differenced to detect & quantify surface motion 	<p>autoRIFT</p> <ul style="list-style-type: none"> • Glacier velocity tracking • Measures surface displacements between two SAR images • Part of NASA MEaSUREs ITS_LIVE project 
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Order an SBAS InSAR Stack in Vertex and Analyze it in OpenSARLab

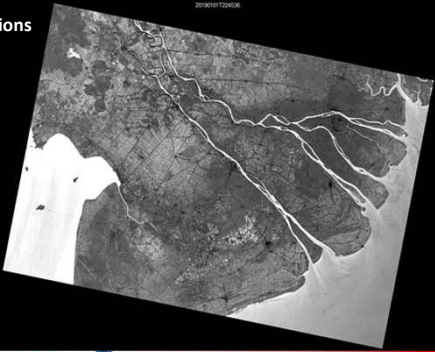
A demo on how to generate an SBAS InSAR Stack in Vertex and then Analyze the data using MintPy in OpenSARLab can be found here:

<https://youtu.be/tq8nZpsWK6k>



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Creating Product Subscriptions over your AOI on HyP3

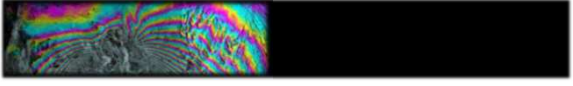


Example of a Sentinel-1 SAR Image Time Series:

- Mekong River Delta
- Time frame: 2019

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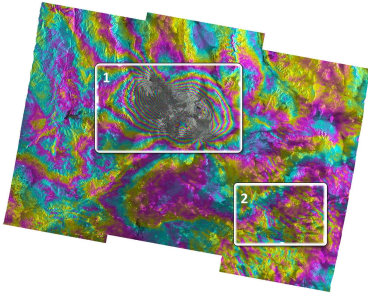
LIMITATIONS OF CONVENTIONAL INSAR



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Think – Pair – Share: Take 1

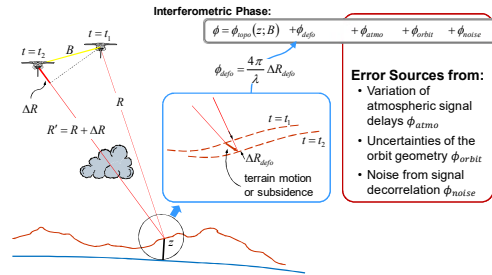
Phase Interpretation: Analyze the phase content of this interferogram



- Activity 1: What do the phase patterns in Area 1 represent?**
 - Displacement phase most likely related to an earthquake
 - Topographic signal
 - Atmospheric delay
- Activity 2: What is the main phase contribution in Area 2?**
 - Displacement phase most likely related to an earthquake
 - Topographic signal
 - Atmospheric delay

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Recap of InSAR Acquisition Scenario



Interferometric Phase:

$$\phi = \phi_{topo}(z; B) + \phi_{defo} + \phi_{atmo} + \phi_{orbit} + \phi_{noise}$$

$$\phi_{defo} = \frac{4\pi}{\lambda} \Delta R_{defo}$$

Error Sources from:

- Variation of atmospheric signal delays ϕ_{atmo}
- Uncertainties of the orbit geometry ϕ_{orbit}
- Noise from signal decorrelation ϕ_{noise}

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The Complete InSAR Phase Equation

- Phase of an interferogram:**

$$\phi = W \{ \phi_{topo} + \phi_{defo} + \phi_{atmo} + \phi_{orbit} + \phi_{noise} \}$$

$$= W \left\{ \frac{4\pi}{\lambda} \frac{B_{\perp}}{R \cdot \sin(\theta)} h + \frac{4\pi}{\lambda} v \cdot \Delta t + \phi_{atmo} + \phi_{orbit} + \phi_{noise} \right\}$$

(W: wrapping operator $\rightarrow \phi: [-\pi, \pi]$)
- Phase of a differential interferogram (after compensation of topography phase):**

$$\Delta\phi = W \left\{ \frac{4\pi}{\lambda} \frac{B_{\perp}}{R \cdot \sin(\theta)} h_{err} + \frac{4\pi}{\lambda} v \cdot \Delta t + \phi_{atmo} + \phi_{orbit} + \phi_{noise} \right\}$$

(where $h_{err} = h_{true} - h_{DEM}$ is a residual topography phase due to errors in the DEM)

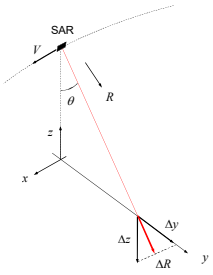
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Limitations and Error Sources of InSAR
 Overview

1. Only sensitive to motion in sensor's line-of-sight → does not provide 3D motion fields
2. Temporal baseline is limited, leading to limited sensitivity to very slow surface motion:
 - Limitation is due to temporal decorrelation, leading to increase of phase noise with time
3. Spatial baseline is limited, limiting number of interferograms that can be formed from a stack of SAR data [advanced topic – not as relevant with sensors such as Sentinel-1 and NISAR]
4. Atmospheric phase patterns may mask signal of interest, limiting sensitivity of InSAR to very small motion (or topography) signals
5. Orbit errors may cause ramp-like phase distortions [usually small]

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Limitations of InSAR:
 1. Only Line-of-Sight Motion Sensitivity



$\Delta R = \Delta y \sin \theta - \Delta z \cos \theta$

for ERS:
 1 fringe (2π) corresponds to

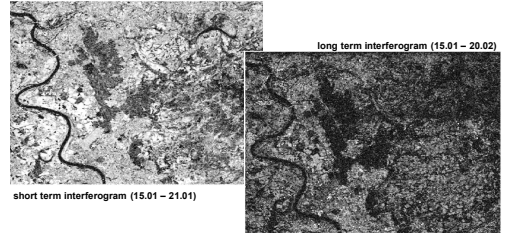
- 2.8 cm in R
- 3.0 cm in z (e.g. subsidence)
- 7.2 cm in y (motion)

! only 1 dimension of 3-d motion accessible
 ! no sensitivity to motion in x direction

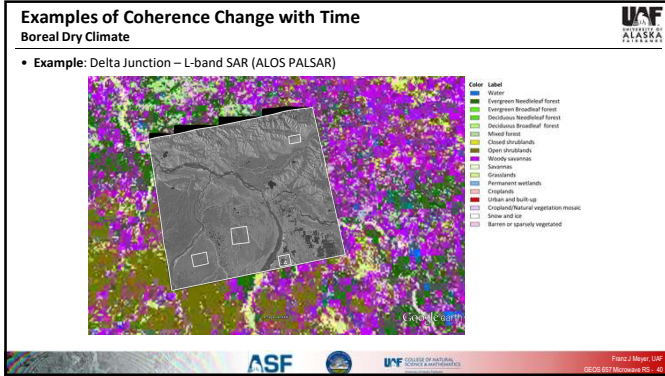
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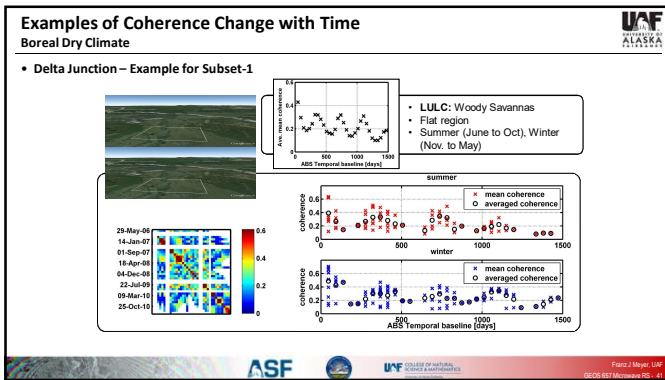
Limitations of InSAR:
 2. Temporal Signal Decorrelation

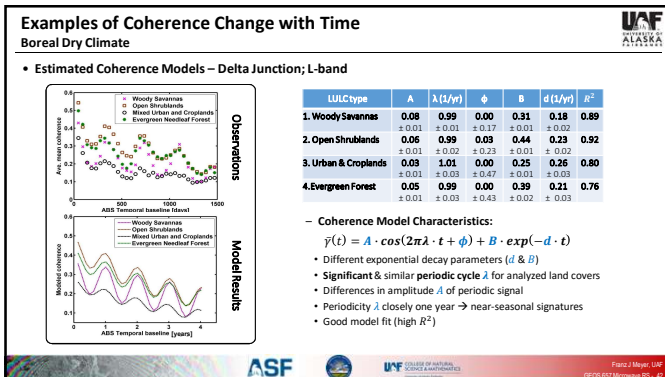
- Changes of the ground scattering properties are main source of signal decorrelation with time
 - Example: Airborne C-band SAR over vegetated environments



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Limitations of InSAR:
2. Temporal Signal Decorrelation

Example: Volcano Monitoring

- Often, caldera of active volcanoes is decorrelated due to constant change (snow and ice melt; ash fall; lahars; lava flows; ...).
- Consequence:**
 - Area with strongest signals inaccessible
 - In modeling, shape of source model as well as horizontal location of source hard to define
 - combination with GPS is advised

Caldera of Westdahl Peak is decorrelated in all interferograms

post-eruption InSAR images (several examples)

1 color cycle = 2.8 cm deformation

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Decorrelation and Fringe Washing

Gorkha Earthquake, Nepal, April 2015

Decorrelation:
 phase essentially random

Fringe Washing
 Large strain: fringes closely spaced, difficult to unwrap
 Can decrease coherence

InSAR Time Series Analysis is mitigating the impact of signal decorrelation

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Limitations of InSAR:
3. Limited Spatial Baselines

Advanced Material

- Long spatial baselines lead to decorrelation of interferograms
- Figure:** Fringe density increases with baseline length!
- At certain baseline length (critical baseline $B_{\perp, crit}$) the fringes become closer than the pixel size:
 - phase jump from pixel to pixel $> 2\pi$
 - phase image becomes random

Short Baseline Interferogram
 Long Baseline Interferogram

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Limitations of InSAR: Advanced Material

How to Calculate the Critical Baseline $B_{\perp,crit}$

- We can calculate the fringe frequency in the interferogram from:

$$\Delta f(\alpha) = -\frac{c \cdot \theta_1}{R \cdot \lambda \cdot t \cdot (\theta - \alpha)} \left[\frac{1}{s} \right] \quad \alpha: \text{terrain slope}$$
- Critical baseline if we have one fringe of phase change per pixel, corresponding to $\frac{2\Delta f(\alpha)}{c} = \frac{1}{\rho_r}$ (with range resolution ρ_r):

$$\frac{1}{\rho_r} = -\frac{2 \cdot B_{\perp,crit}}{R \cdot \lambda \cdot \tan(\theta - \alpha)} \Rightarrow B_{\perp,crit} = \left| \frac{R \cdot \lambda \cdot \tan(\theta - \alpha)}{2 \cdot \rho_r} \right| [m]$$
- Dependencies:
 - Higher range resolution (smaller ρ_r) $\rightarrow B_{\perp,crit}$ increases
 - Steeper slopes (larger α) $\rightarrow B_{\perp,crit}$ decreases
 - Longer wavelength $\lambda \rightarrow B_{\perp,crit}$ increases

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Critical Baseline $B_{\perp,crit}$ for Various Sensors Advanced Material

- Assumptions:
 - Surface slope $\alpha = 0^\circ$
 - Incidence angle $\theta = 25^\circ$
 - Range to ground $R \approx 800,000m$
- $B_{\perp,crit}$ for various sensor systems @ $\theta = 25^\circ$:

ERS-1/2 & Envisat Stripmap mode:	$B_{\perp,crit} \approx 1,100m$
ALOS PALSAR:	$B_{\perp,crit} \approx 8,500m$
• FBS mode:	$B_{\perp,crit} \approx 4,250m$
• FBD & PLR mode:	$B_{\perp,crit} \approx 4,250m$
TerraSAR-X Stripmap mode:	$B_{\perp,crit} \approx 5,800m$

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Limitations of InSAR: Advanced Material

4. Atmospheric Distortions

- Atmospheric propagation influence:
 - Masks deformation signals in InSAR
 - Increases data requirements and latency times for InSAR analyses

What we would like

What we often get

-3 cmly subsidence

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Limitations of InSAR:
4. Atmospheric Distortions

Coquimbo Earthquake, Chile, April 2018, M6

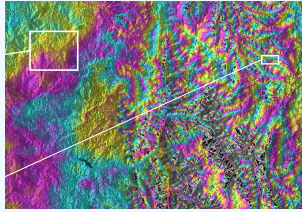
Turbulent Troposphere:
 Horizontal variation in refractive index, esp. in water vapor
 Variable across spatial scales
 Can mask small earthquake

Correction:

- Regional scale: weather model, GPS, etc.
- Small scale: filtering in space and time

Stratified troposphere
 Signal path decreases with elevation
 Phase correlated with topography

Correction: weather models, data-driven

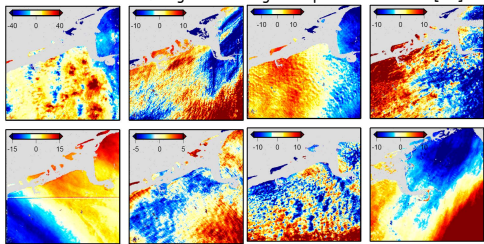


Sentinel-1 Interferogram
 Courtesy SARVIEWS, ESA

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Examples of Turbulent Atmospheric Phase Distortions in InSAR Data

• 8 100x100 km interferograms showing atmospheric distortions [cm]

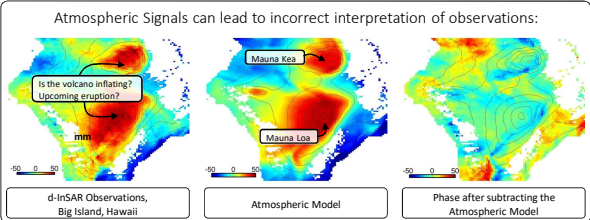


Courtesy: R. Hansen, TU Delft

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Example and Impact of Stratified Atmospheric Phase Distortions

Atmospheric Signals can lead to incorrect interpretation of observations:



Is the volcano inflating?
 Upcoming eruption?

Mauna Kea
 Mauna Loa

d-InSAR Observations, Big Island, Hawaii
 Atmospheric Model
 Phase after subtracting the Atmospheric Model

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Limitations of InSAR:

4. Atmospheric Distortions

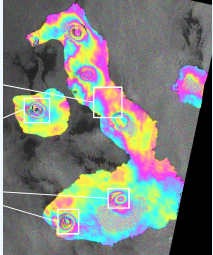
Isabela Island, Galapagos, Ecuador, September 2019, M6

Troposphere or volcanic origin?

Stratified troposphere or volcanic origin?
How could you quantify the relative contributions?

InSAR Time Series Analysis is mitigating the impact of atmospheric phase delay.

Sentinel-1 Interferogram
Courtesy: SARVINSW, ESA



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What's Next?

• This is what awaits next:

- Thursday March 30: Lab on Interpreting InSAR data
- Tuesday April 04: Polarimetric SAR

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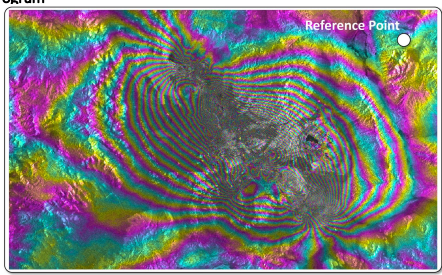
Think - Pair - Share: Take 2

Quantify What's in the Interferogram

• **Activity:** Based on the interferogram and following the instructions above, what is the amount of line-of-sight deformation contained in the image?

- a) 0-25 cm
- b) 60-80 cm
- c) 100-150 cm

Reference Point



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